Department of Theoretical Geophysics & Mantle Dynamics University of Utrecht, The Netherlands

Computational Geodynamics

FDM & the Stokes equation

Cedric Thieulot c.thieulot@uu.nl

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Introduction

Numerical solutions of the momentum and continuity equations

Results

C. Thieulot | Introduction to FDM





Mass conservation equation for incompressible fluids:

 ${oldsymbol
abla}\cdot {oldsymbol v}=0$

Momentum conservation equation

 $-\boldsymbol{\nabla}\boldsymbol{\rho} + \boldsymbol{\nabla} \cdot \boldsymbol{s} = \rho \boldsymbol{g}$ with $\boldsymbol{s} = 2\mu \dot{\boldsymbol{\epsilon}} = \mu (\boldsymbol{\nabla} \boldsymbol{v} + \nabla \boldsymbol{v}^{T})$

$$-\frac{\partial p}{\partial x} + \frac{\partial s_{xx}}{\partial x} + \frac{\partial s_{xy}}{\partial y} = \rho g_x$$
$$-\frac{\partial p}{\partial y} + \frac{\partial s_{xy}}{\partial x} + \frac{\partial s_{yy}}{\partial y} = \rho g_y$$



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or,

$$-\frac{\partial p}{\partial x} + 2\mu \frac{\partial \dot{\epsilon}_{xx}}{\partial x} + 2\mu \frac{\partial \dot{\epsilon}_{xy}}{\partial y} = \rho g_x$$
$$-\frac{\partial p}{\partial y} + 2\mu \frac{\partial \dot{\epsilon}_{xy}}{\partial x} + 2\mu \frac{\partial \dot{\epsilon}_{yy}}{\partial y} = \rho g_y$$



In 2D, we first explicitely write

$$\dot{\epsilon}_{xx} = \frac{\partial u}{\partial x}$$
 $\dot{\epsilon}_{xy} = \frac{1}{2} \left(\frac{\partial u}{\partial y} + \frac{\partial v}{\partial x} \right)$ $\dot{\epsilon}_{yy} = \frac{\partial v}{\partial y}$

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Then

$$-\frac{\partial \rho}{\partial x} + 2\mu \frac{\partial \dot{\epsilon}_{xx}}{\partial x} + 2\mu \frac{\partial \dot{\epsilon}_{xy}}{\partial y} = \rho g_x$$

becomes

$$-\frac{\partial p}{\partial x} + 2\mu \frac{\partial^2 u}{\partial x^2} + \mu \left(\frac{\partial^2 u}{\partial y^2} + \frac{\partial^2 v}{\partial x \partial y}\right) = \rho g_x$$



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Using $\nabla \cdot \boldsymbol{v} = 0$ to obtain $\frac{\partial u}{\partial x} = -\frac{\partial v}{\partial y}$ we arrive at

$$-\frac{\partial p}{\partial x} + \mu \left(\frac{\partial^2 u}{\partial x^2} + \frac{\partial^2 u}{\partial y^2}\right) = \rho g_x$$



These are then the two coupled equations we need to solve

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along with

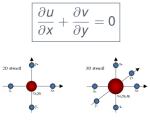
$$\boxed{\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} = 0}$$

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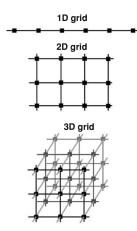
along with



2D/3D stencils for Laplacian

Num. solutions of the Stokes eqs Grids & stencils





From Gerya, 2010

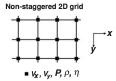


Fig. 7.5 Example of a non-staggered 2D numerical grid.

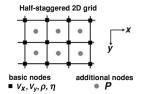


Fig. 7.6 Example of a half-staggered 2D numerical grid.



 Fully staggered grids are applied in 2D and consist of a combination of several types of nodal points having different geometrical positions.



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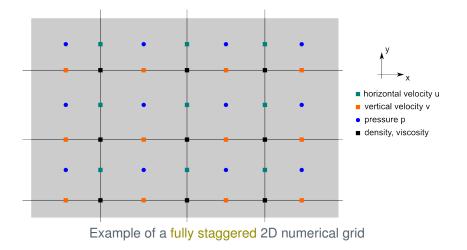


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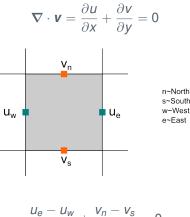


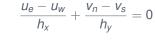
- Fully staggered grids are applied in 2D and consist of a combination of several types of nodal points having different geometrical positions.
- Different variables are then defined at different nodal points.
- Different equations are also formulated at different nodal points.
- Despite the apparent geometrical complexity, fully staggered grids are the most convenient choice for thermomechanical numerical problems with variable viscosity when finite differences are used for solving the Stokes and temperature equations.

Num. solutions of the Stokes eqs Grids & stencils



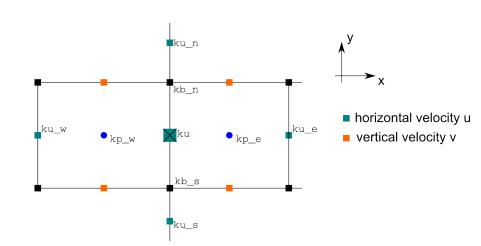
For incompressible flow:







Num. solutions of the Stokes eqs 2D staggered grid, x-direction

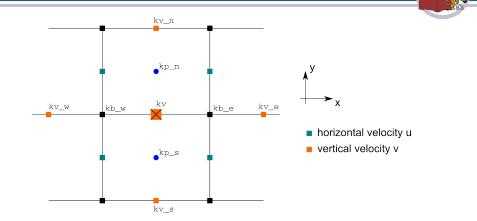


Num. solutions of the Stokes eqs 2D staggered grid, x-direction

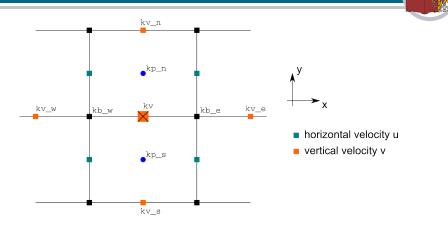


$$\begin{aligned} -\frac{\partial p}{\partial x}\Big|_{ku} &= -\frac{p_{kp_e} - p_{kp_w}}{h_x} \\ \mu \frac{\partial^2 u}{\partial x^2}\Big|_{ku} &= \mu \frac{u_{ku_e} - 2u_{ku} + u_{ku_w}}{h_x^2} \\ \mu \frac{\partial^2 u}{\partial y^2}\Big|_{ku} &= \mu \frac{u_{ku_n} - 2u_{ku} + u_{ku_w}}{h_y^2} \\ \rho g_x\Big|_{ku} &= \frac{\rho_{kb_n} + \rho_{kb_w}}{2} g_x \end{aligned}$$

2D staggered grid, y-direction

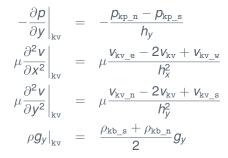


2D staggered grid, y-direction



- Similar expressions can be written for the *z*-component of the momentum equation.
- Not difficult, only careful bookkeeping







Mechanical boundary conditions depend on the type of numerical problem which is studied.

The following boundary conditions are often used in geomodelling:

- ► free slip
- no slip
- free surface
- fast erosion
- ▶ infinity-like (external free slip, external no slip, Winkler basement)
- prescribed velocity (moving boundary)
- periodic
- combined conditions

- 15
- free slip b.c.: the normal velocity component on the boundary is zero and the two other components do not change across the boundary (this condition also implies zero shear strain rates and stresses along the boundary). For example, for the boundary orthogonal to the x axis, the free slip condition is formulated as follows

$$u = 0 \qquad \qquad \frac{\partial v}{\partial x} = \frac{\partial w}{\partial x} = 0$$

A no slip condition requires all velocity components on the boundary to be zero, i.e.

$$u = v = w = 0$$

 A free surface condition requires both shear and normal stresses at the boundary to be zero

$$s_{ij}=0$$



The prescribed velocity condition implies non-zero velocity at a model boundary. When velocity is prescribed orthogonal to the boundary (inward/outward flow), then a compensating outward/inward velocity should be prescribed on the other model boundary(ies) in order to insure mass conservation in the model.

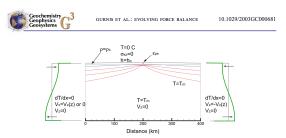


Figure 5. Model setup showing boundary and initial conditions. The initial temperature field is shown with 300°C isotherms in red for Case 15 (as an example). The velocity boundary conditions are shown schematically off to the left and the right of the domain. The depth scale is the same as the horizontal.



- Periodic boundary conditions are typically established for paired parallel lateral boundaries of a model and prescribe that all material properties as well as pressure and velocity fields at both sides of each boundary are identical. From a physical point of view, this implies that these two boundaries are open and that flow leaving the model through one boundary immediately re-enters through the opposite side. This condition is often used in mantle convection modelling to simulate part of a spherical/cylindrical shell with a convecting mantle (or mimic it, in Cartesian coordinates).
- Combined conditions represent a mixture between several types of boundary conditions.



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- Boundary conditions can also be applied inside the model.



implementation free and no slip for staggered grid p94



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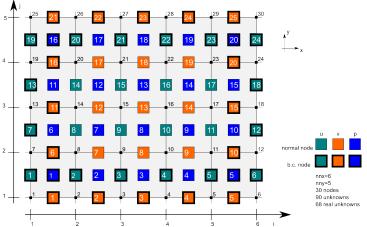
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- This is a somewhat boring subject but it is extremely important to understand it properly. (Remember, 90% of the bugs in your code are made with the indexing).



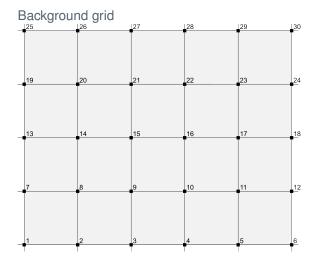
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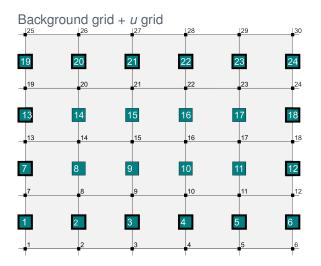
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- One of the (optimal?) ways of numbering is illustrated hereafter:



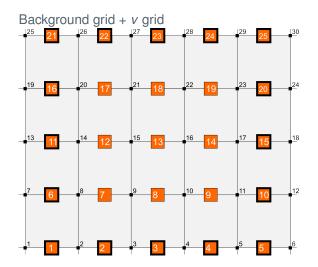




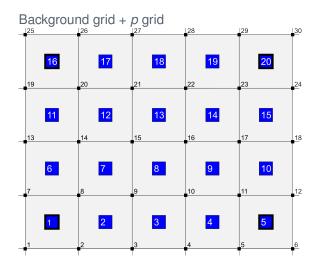










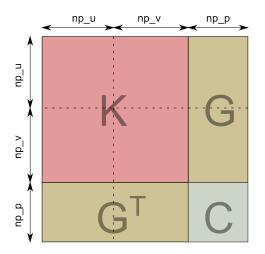




The discretisation of the Stokes equations by means of the FDM or FEM yield a discrete system of equations which takes the form:

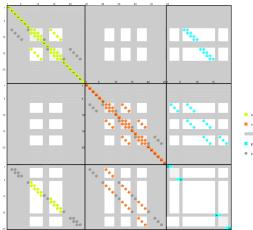
$$\left(\begin{array}{cc} K & G \\ G^T & 0 \end{array}\right)$$





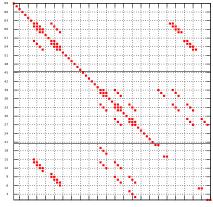
Num. solutions of the Stokes eqs Detailed matrix structure





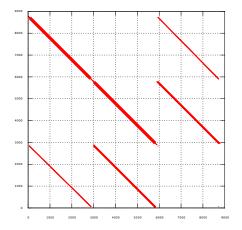
- v be
- zeros due to bo
- pressure comers
- zeroed values because of bc





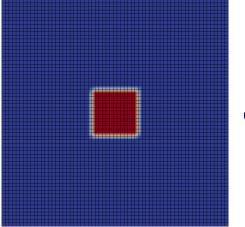
6 9 12 15 18 21 24 27 30 33 36 39 42 45 48 51 54 57 60 63 66 69

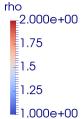






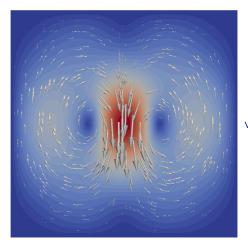












velocity 2.740e-01 0.206 0.137 0.069 0.000e+00



